



CO₂QUEST

Potential Impact of Selected Impurities on Geochemistry Related to CO₂ Storage

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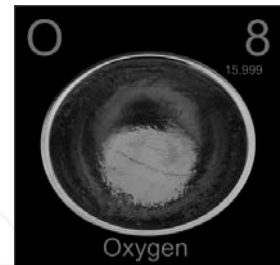
International Forum on Recent Development of CCS Implementation

26 – 27 March 2015, Ledra Hotel, Athens, Greece

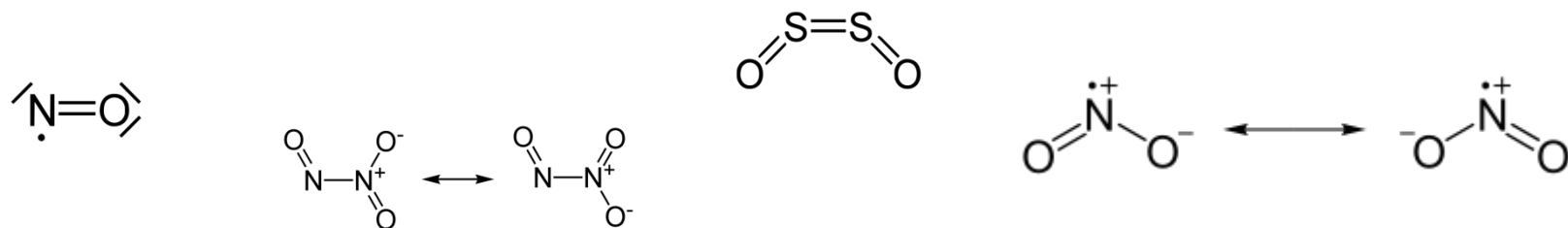


impact of non-condensable gases

- changing physical properties of CO₂
mobility, density, wettability
→ **injectivity**
- use of pore space
→ smaller storage volume for CO₂

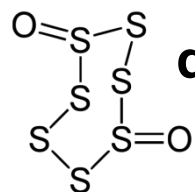


impact of impurities – chemistry

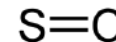
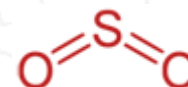
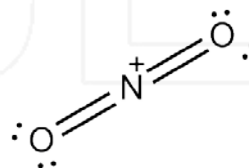
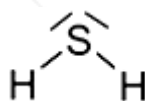
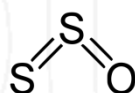
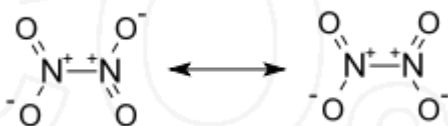
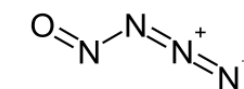
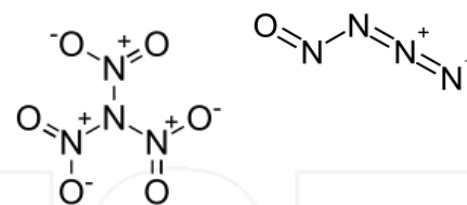
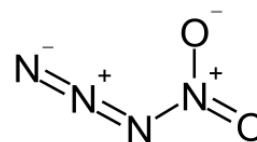
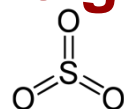


impact of chemically active gases – H_2S , NO_x , SO_x

- acidification around the well
dissolution of carbonates



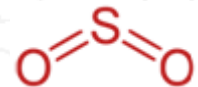
- acidification in greater distances
precipitation of secondary minerals





impact of chemically active gases – H_2S , NO_x , SO_x

- **acidification around the well**
dissolution of carbonates
- **acidification in greater distances**
precipitation of secondary minerals

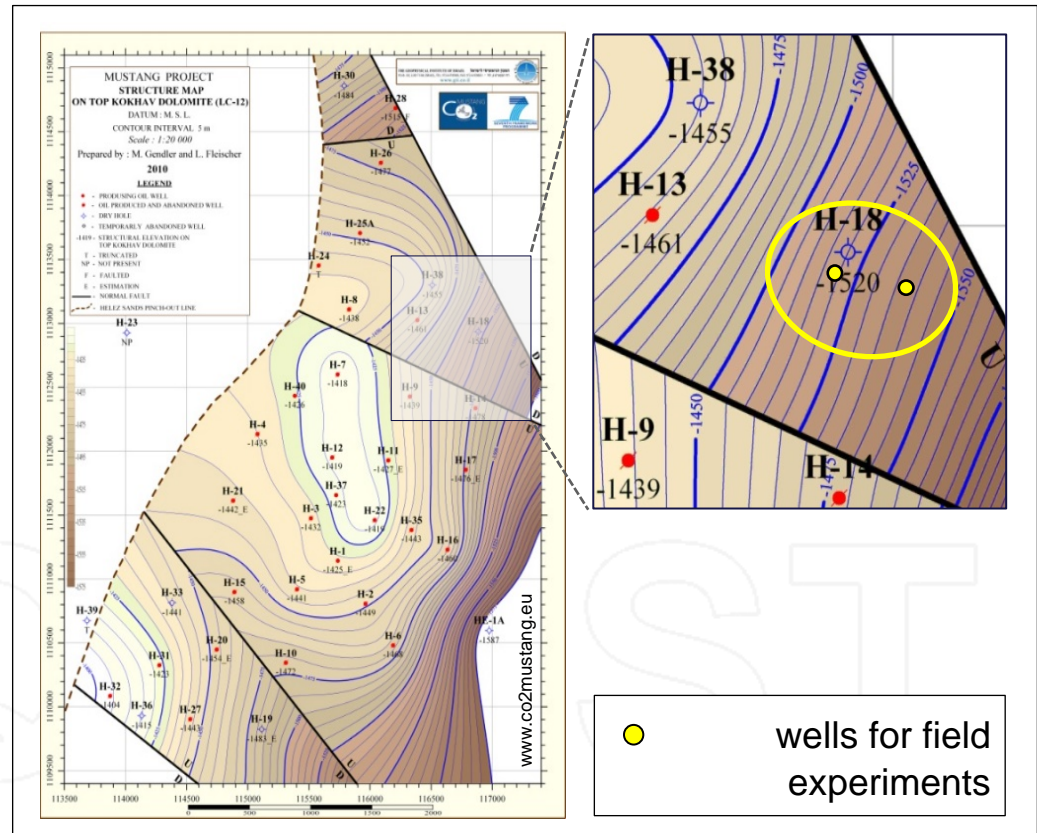


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field scale experiment

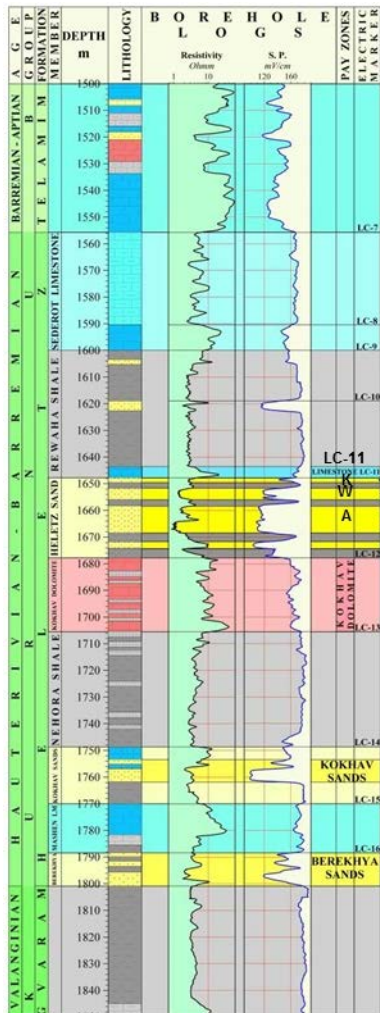
Southern Mediterranean
Coastal Plain of Israel

saline water on
edges of
depleted oil field





reservoir structure



caprock, shales + marls
limestone layer LC11
3 sand reservoir layers
K, W, A

anticlinal structure

thickness of target horizon approx. 20 m

3 conductive, Lower Cretaceous sandstones K, W, A
with cumulative thickness of approx. 10 m

QUEST



parameter	units	value	error
→ temperature	°C	66	± 1
→ pressure	MPa	14.7	± 1
salinity	%-mass	0.055	± 0.005
→ porosity	%	20	± 5
→ horizontal permeability	mD	700	$\times/\div 10$
vertical permeability	mD	100	$\times/\div 10$
rock grain density	kg/m ³	2870	± 75
rock bulk density	kg/m ³	2300	± 300
wet thermal conductivity	W/m/K	4.0	± 0.8
dry thermal conductivity	W/m/K	2.6	± 0.6
specific heat	J/kg/K	920	± 100
thickness of target horizon	m	11, with 2 layers: 2 m + 9 m	± 4 to 6

core samples
lab measurements

[Benson et al., 2013]
[Bensabat et al., 2013]
[Cermak + Rybach, 1992]
[Edlmann, 2013]
[Niemi et al., 2014]
[Pruess et al., 2012]
[Shtivelman, 2010]
[Vosteen + Schellschmidt, 2003]



minerals (I)

primary minerals		volume fraction [%]	secondary minerals	
carbonates				
→ ankerite	e.g. $\text{CaFe}(\text{CO}_3)_2$	3.7	calcite	CaCO_3
			siderite	FeCO_3
feldspar				
K-feldspar	KAlSi_3O_8	12		
albite	$\text{NaAlSi}_3\text{O}_8$	2.5		
clay minerals				
illite	e.g. $\text{K}_{0.85}\text{Al}_{2.85}\text{Si}_{3.15}\text{O}_{10}(\text{OH})_2$	3.9		
chlorite	$(\text{Mg,Fe})_3(\text{Si,Al})_4\text{O}_{10}$	1.4		
kaolinite	$\text{Al}_2\text{Si}_2\text{O}_5(\text{OH})_4$	3.2		



minerals (II)

primary minerals		volume fraction [%]	secondary minerals	
sulfur minerals				
→ anhydrite	CaSO_4	0.4		
→ pyrite	FeS_2	2.1		
sandstone				
quartz	SiO_2	69.3		
iron minerals				
goethite	FeOOH	1.5	hematite	Fe_2O_3

- thermodynamic data – EQ3/6 database and [Wolery, 1992]
- kinetic data – [Palandri and Kharaka, 2004]
[Zhang et al, 2011], [Xu et al., 2006]



1D model

- radial symmetry
- injection of CO_2 with 1 % SO_2

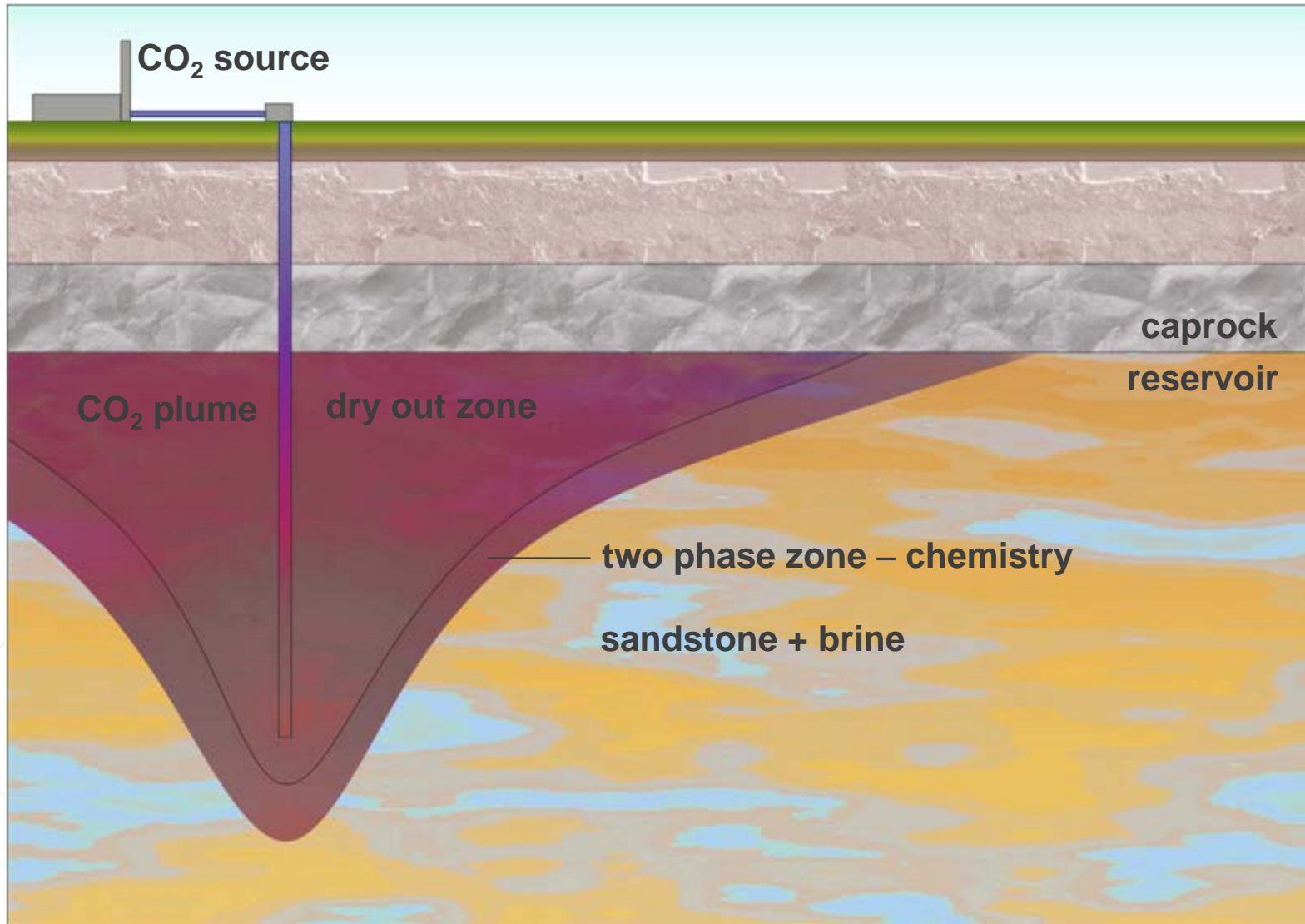
scope

- different time scales
- transport vs. aqueous chemistry
- mineral reaction

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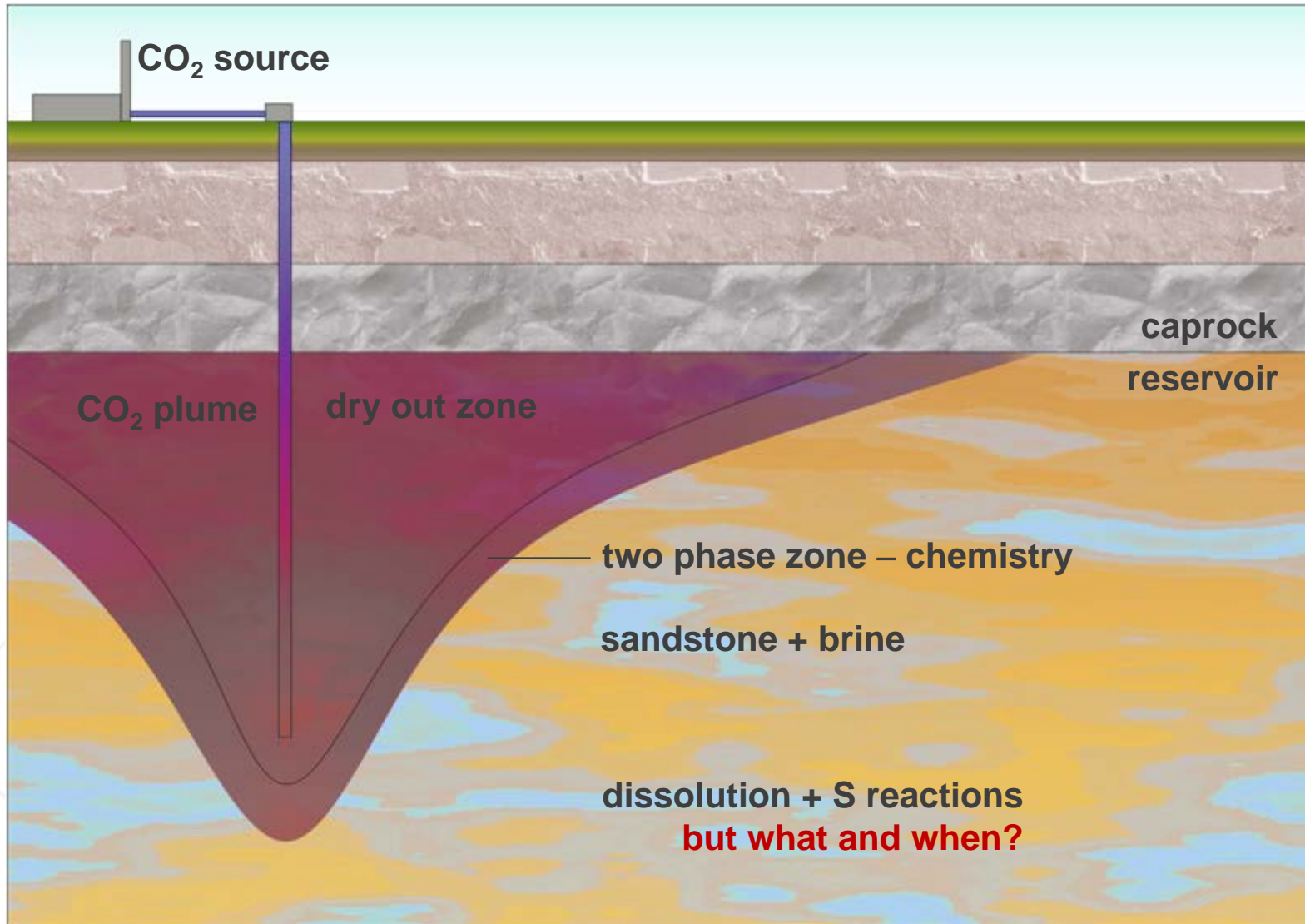
reservoir

area of interest



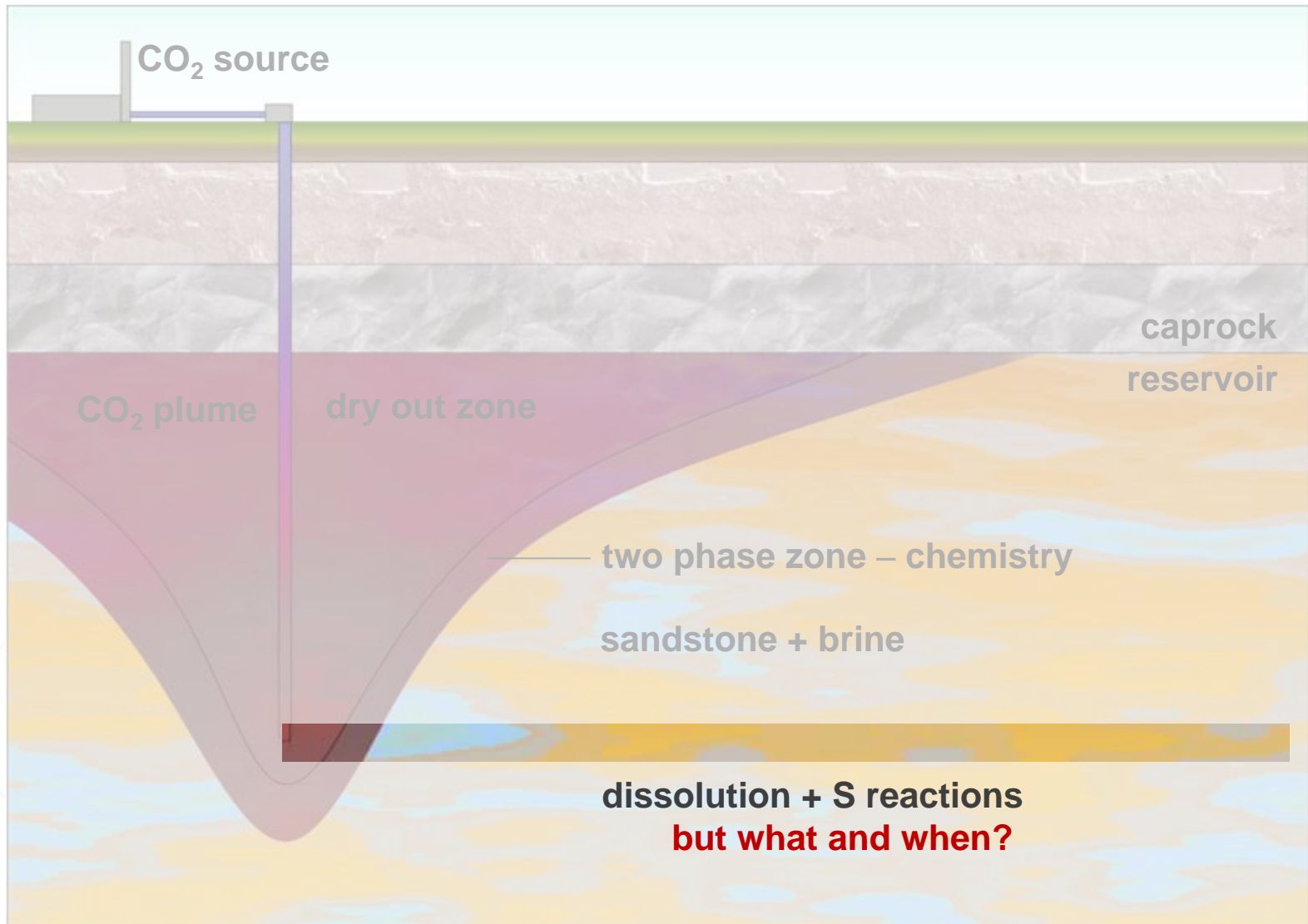
reservoir

area of interest



reservoir

area of interest





1D model, radial

- horizontal 30 km
of chemical interest 100 m to 1500 m
- vertical 11 m

cells

- total 500
- cell size 0.112 m to 2545 m

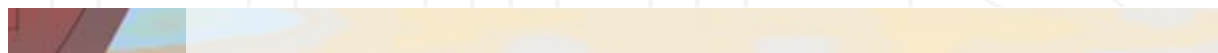
injection

CO₂ with 1 % SO₂, 9.0 kg/s
simulation period 10 a



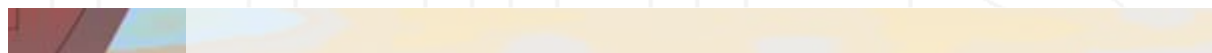
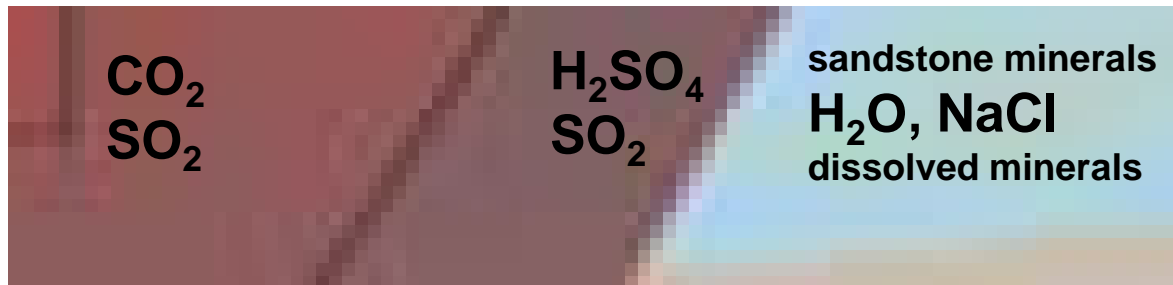


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chemically of interest



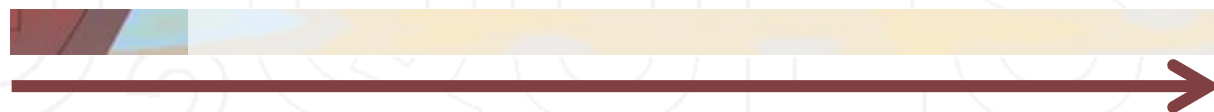
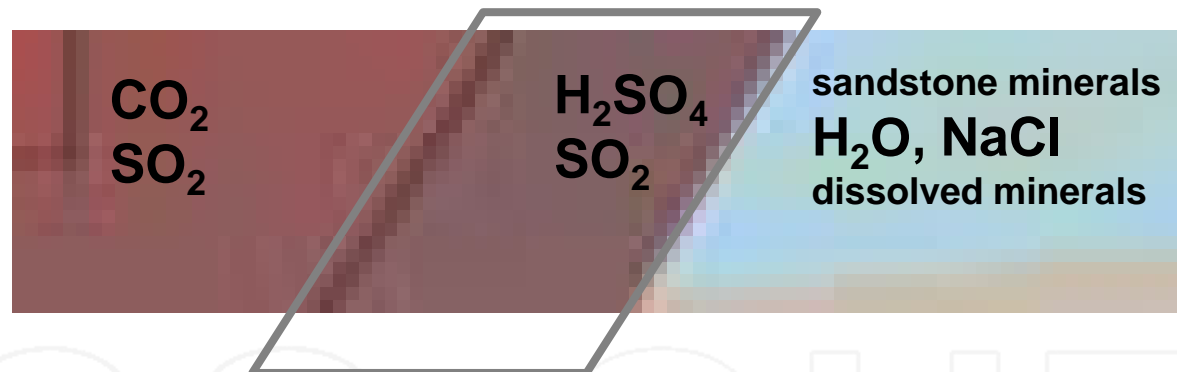
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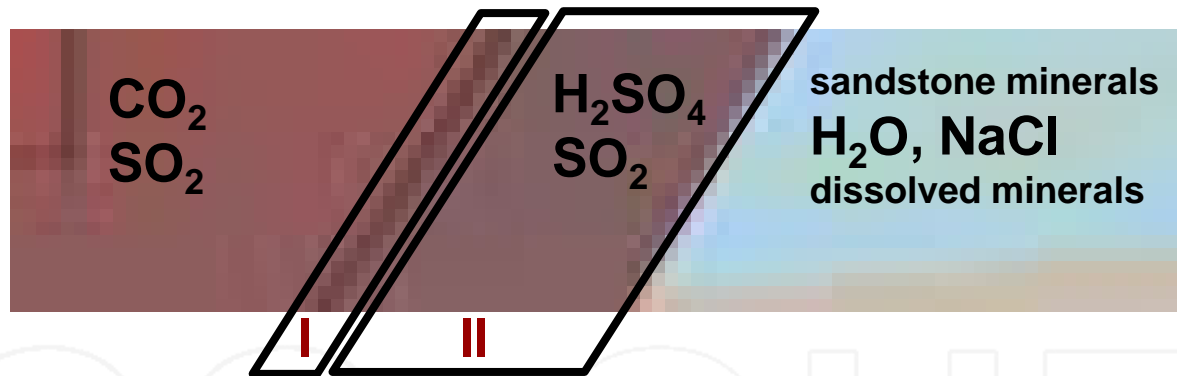
chemically of interest

dry out zone

two phase zone pressure zone



transport



CCO₂QUEST



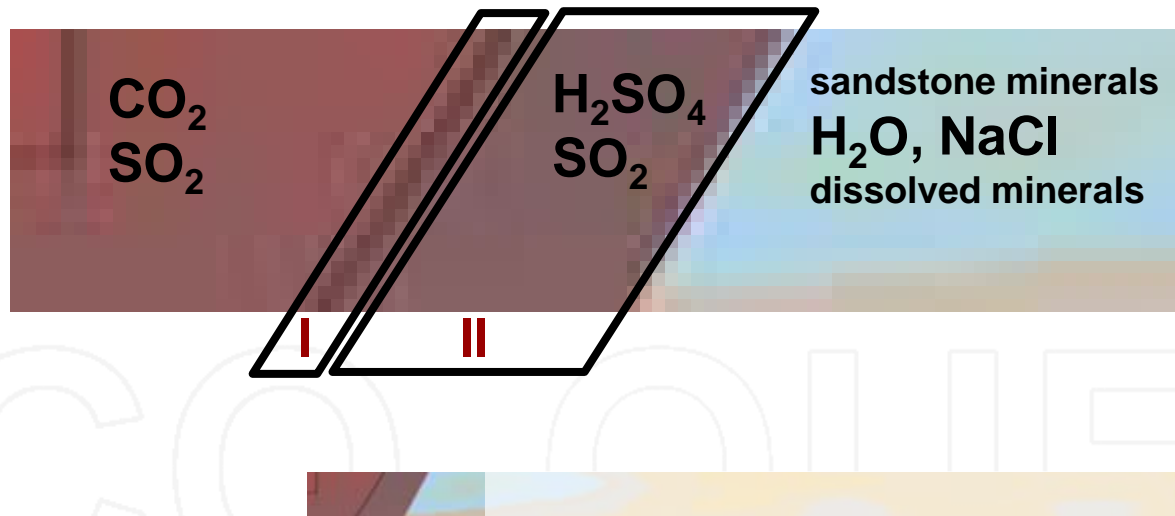
in two phase zone distinguish reactive zones I and II

(I)

- closer to injection point
- high dissolution rate
- narrow zone

(II)

- farther away
- slow, only small chromatographic dissociation
- SO_2 distribution in broad zone
- low concentration





injection

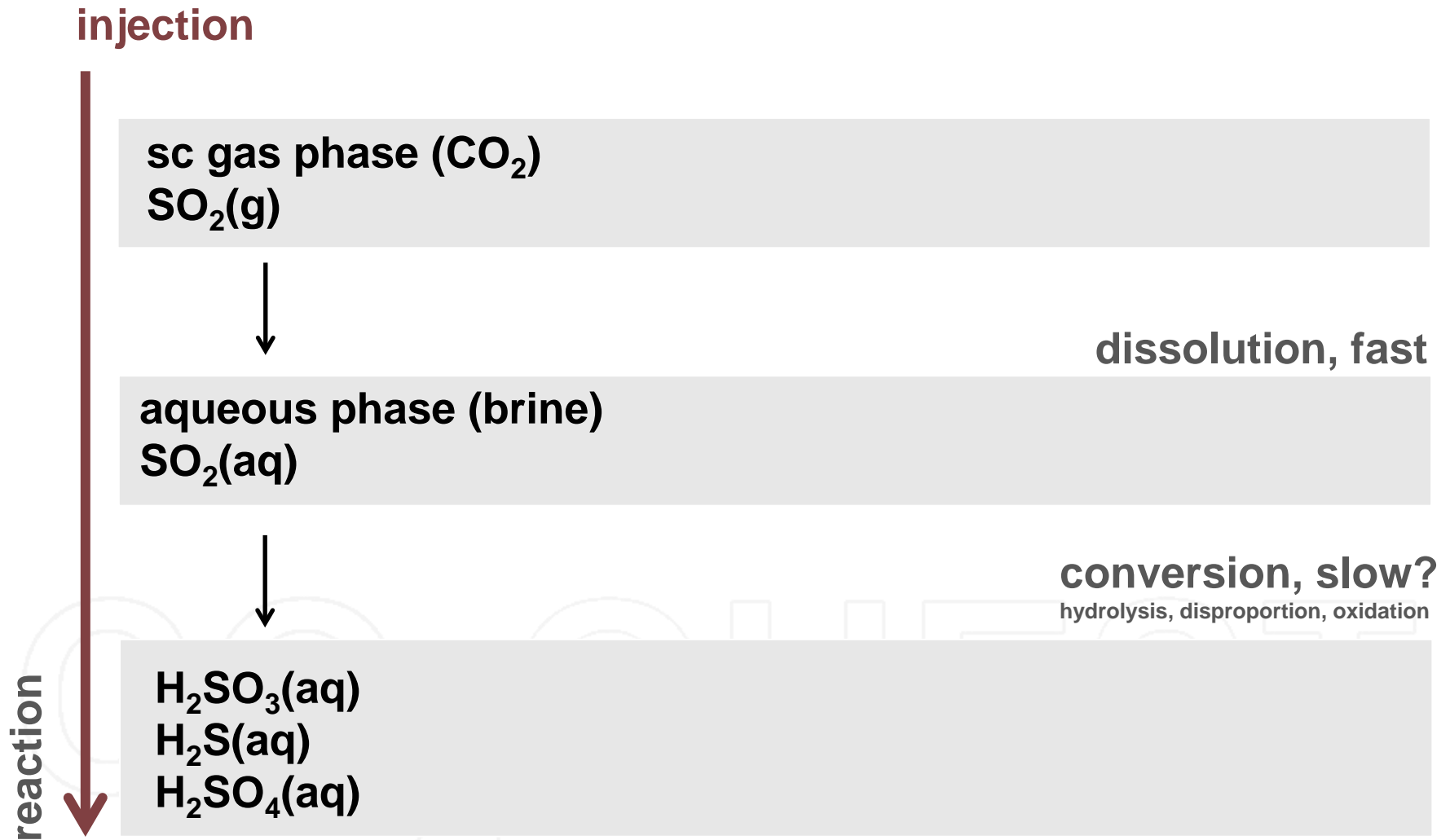
sc gas phase (CO_2)
 $\text{SO}_2(\text{g})$

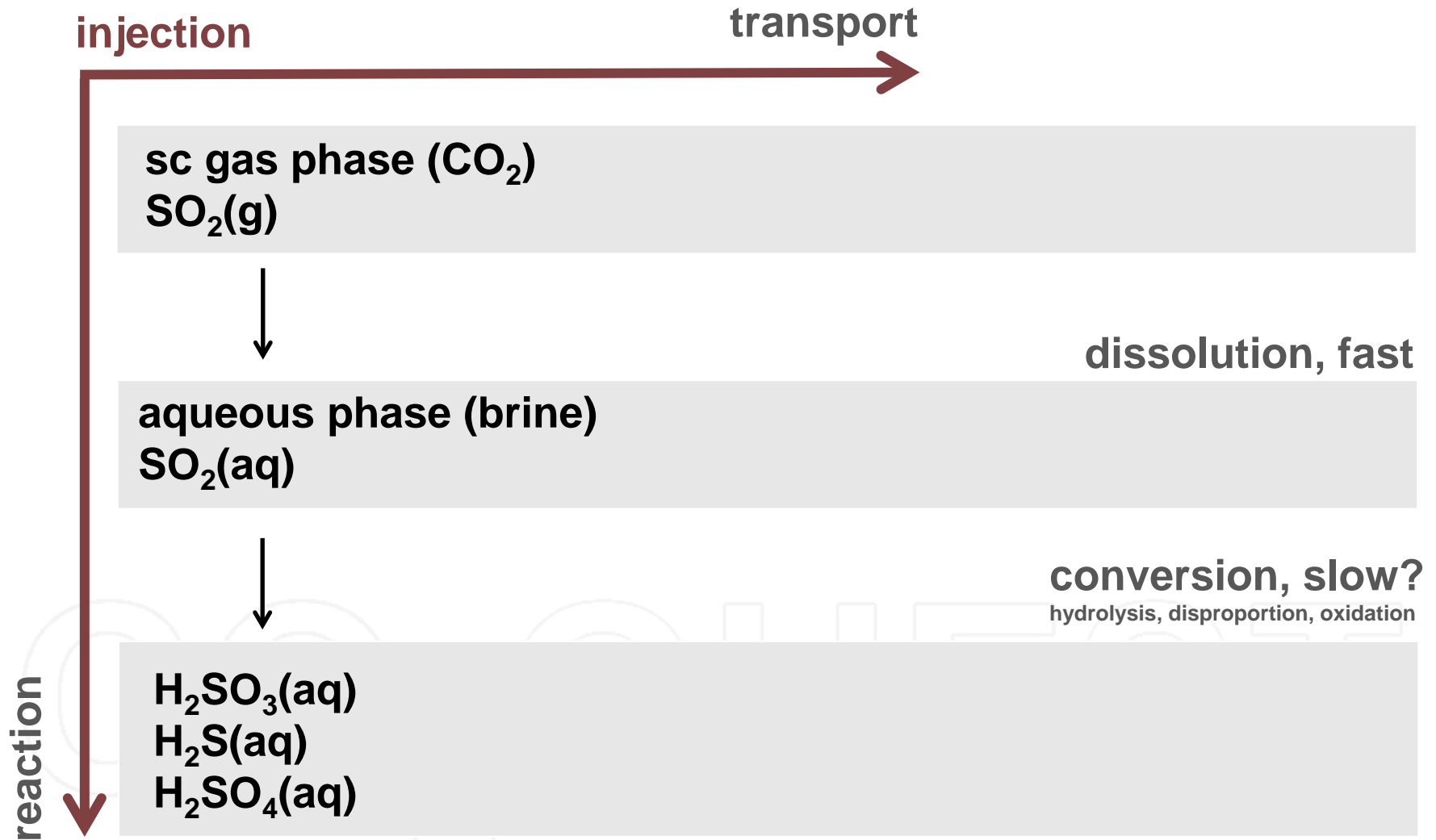


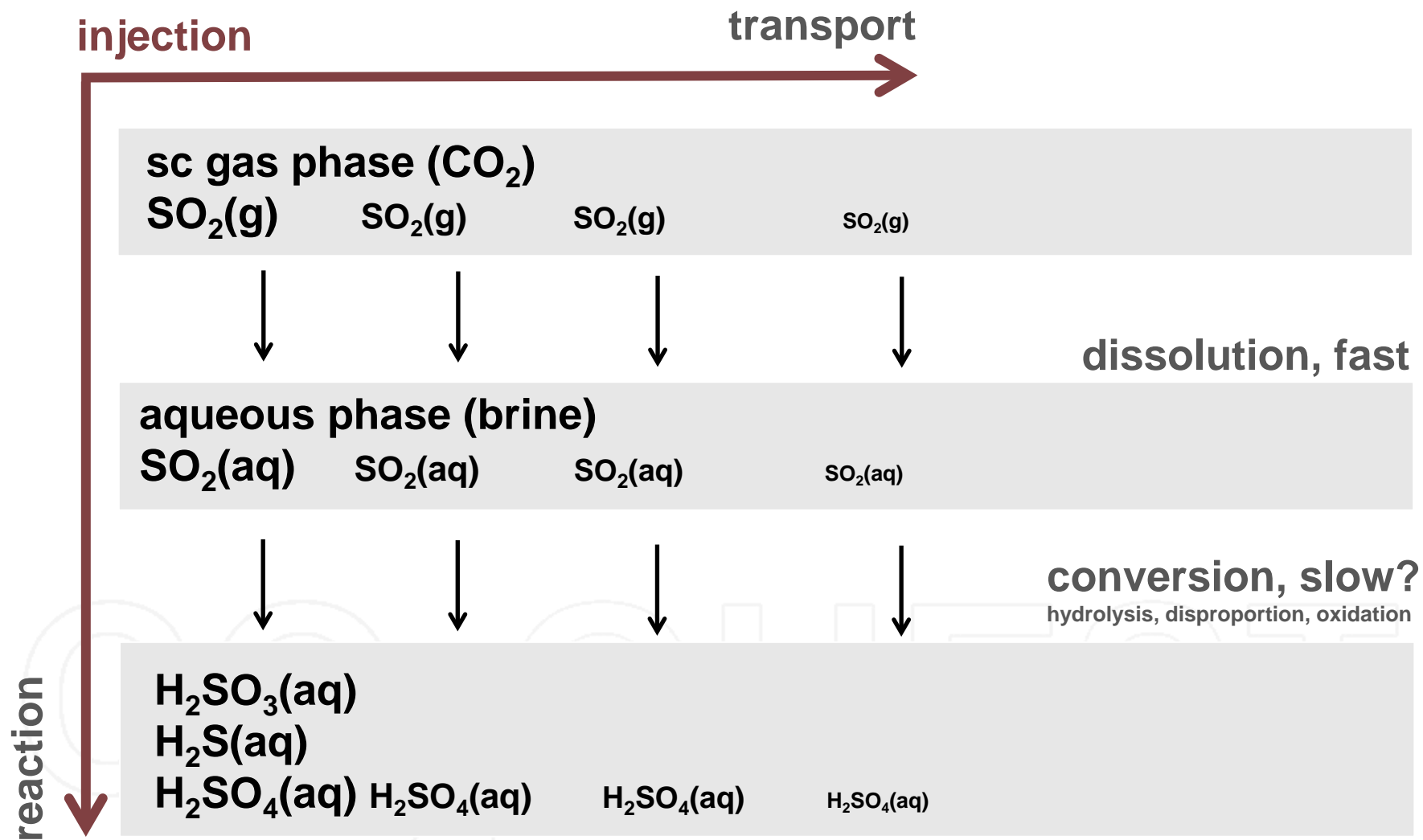
aqueous phase (brine)
 $\text{SO}_2(\text{aq})$



$\text{H}_2\text{SO}_3(\text{aq})$
 $\text{H}_2\text{S}(\text{aq})$
 $\text{H}_2\text{SO}_4(\text{aq})$









SO₂ dissolution and oxidation

main problem

→ **chemical conversion rates of SO₂
under geological storage conditions
are controversial and/or unknown**

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transport vs. SO₂ oxidation rate

$$r_{\text{transp}} = \frac{dn}{dt \cdot dV_{\text{aq}}} = \frac{c_{\text{gas}} \cdot S_g}{dt \cdot (1 - S_g)}$$

0.02 mol·L⁻¹·s⁻¹ to 0.07 mol·L⁻¹·s⁻¹

$$r_{\text{oxidation}} = k_{(T=66^\circ\text{C}, p=14.7\text{MPa})} \cdot a(\text{SO}_{2(aq)})$$

10⁻¹⁰ mol·L⁻¹·s⁻¹ to 10⁻¹ mol·L⁻¹·s⁻¹

rate studies at atmospheric conditions — but geological conditions

higher temperature → increase of reaction rate

higher salinity → decrease of reaction rate

equilibrium speciation — infinitely fast conversion

kinetically delayed speciation — slow conversion

applied value for oxidation: 5·10⁻⁷ mol·L⁻¹·s⁻¹



Transport Of Unsaturated Groundwater + Heat

THMC numerical simulation

multi-dimensional

coupled transport of water, vapor, non-condensable gas, and heat

porous and fractured media

LBNL Berkeley

finite differences

→ **injection of dissolved SO_2 in gas phase**

TOUGHREACT 3.0-OMP + ECO2N

simulation of multiphase flow of pure gases

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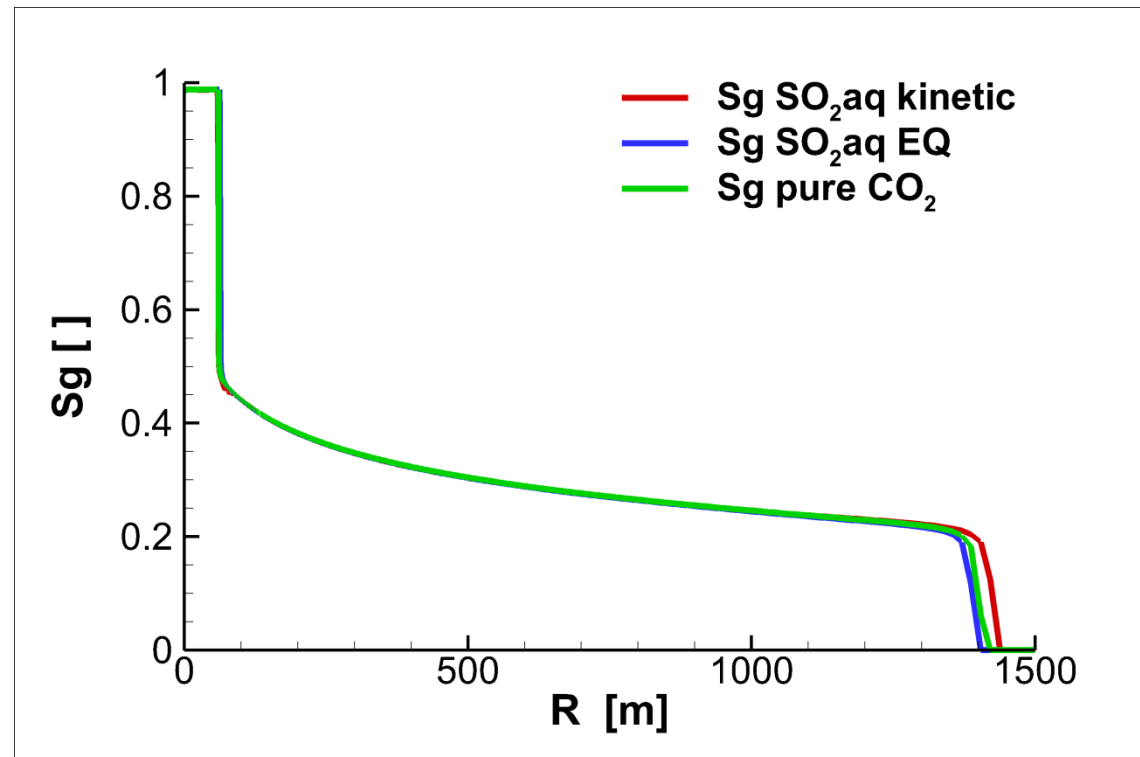


gas saturation

$t = 10 \text{ a}$

compare
pure CO_2 with
 $\text{CO}_2 + \text{SO}_2$
kinetic and
equilibrium

no difference

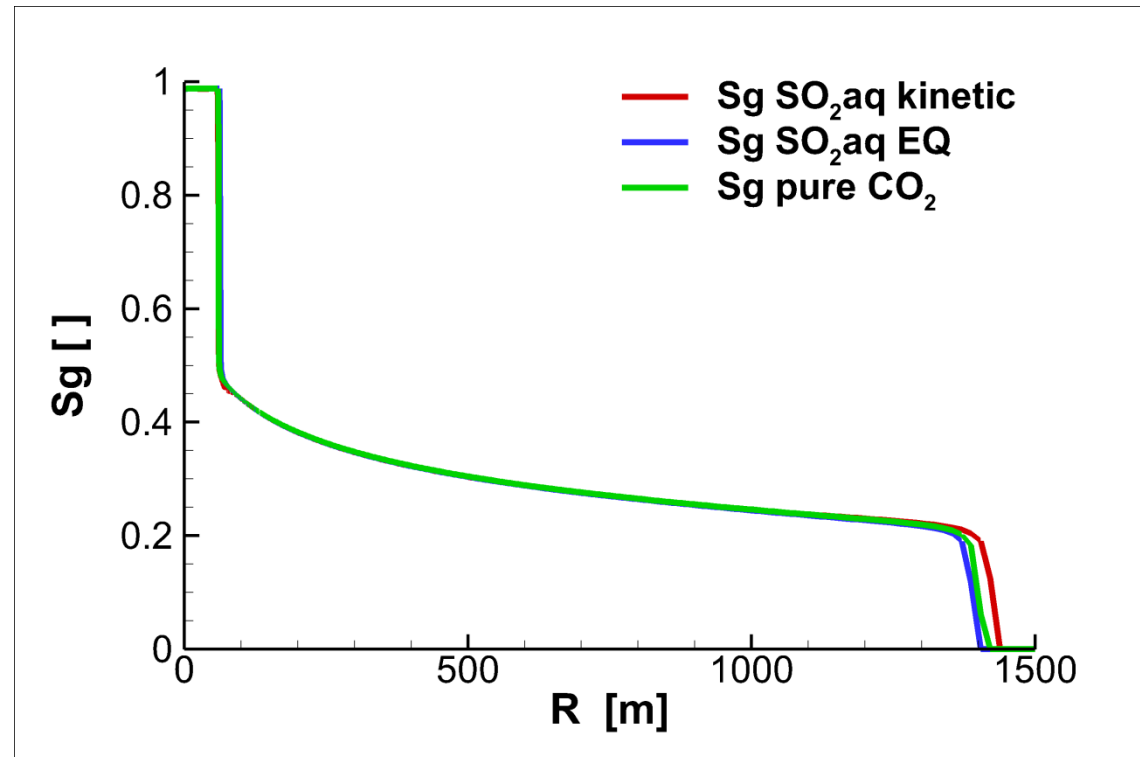




gas saturation

$t = 10 \text{ a}$

no difference



dry out zone →

< 100 m

two phase zone

> 100 m

pressure zone

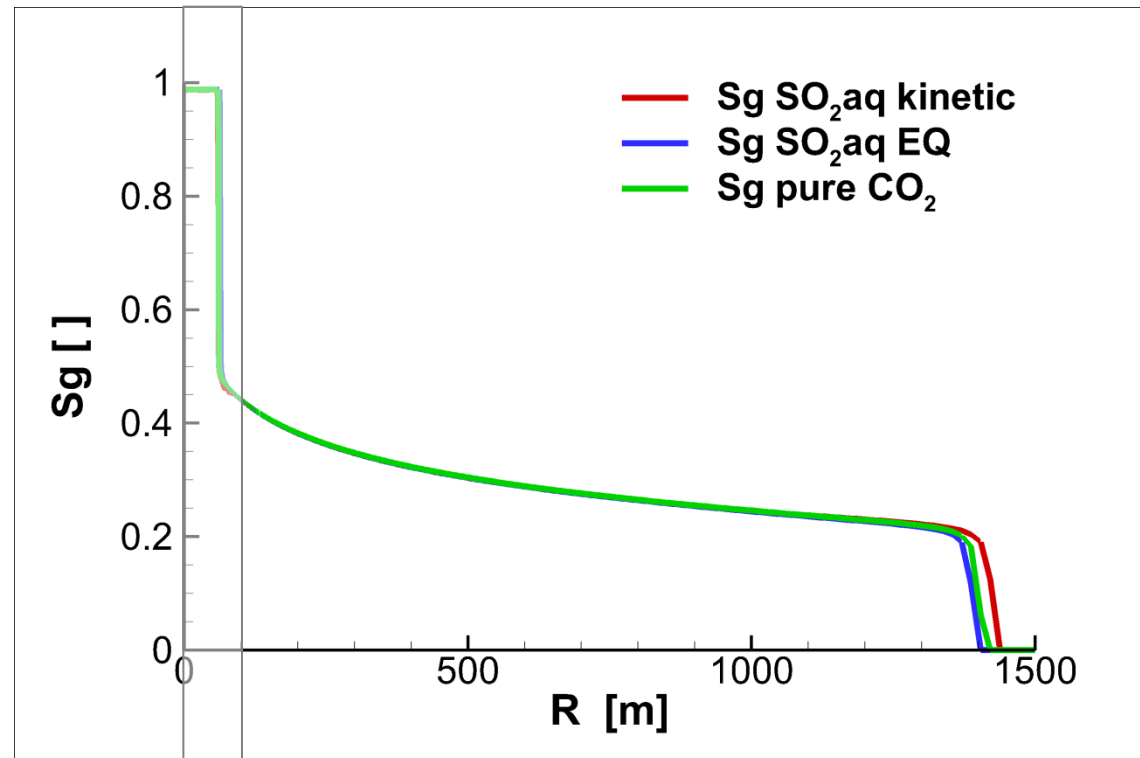
> 1400 m



gas saturation

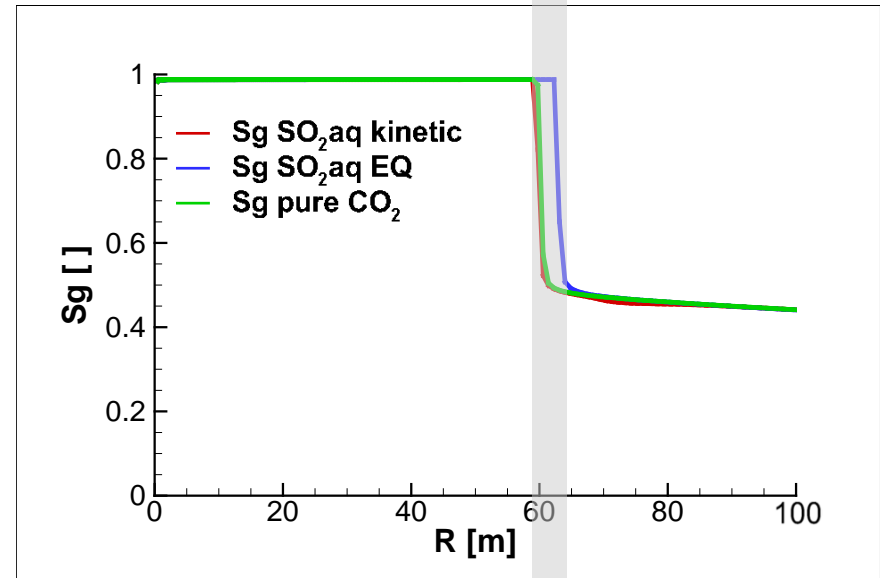
$t = 10$ a

no difference





saturation



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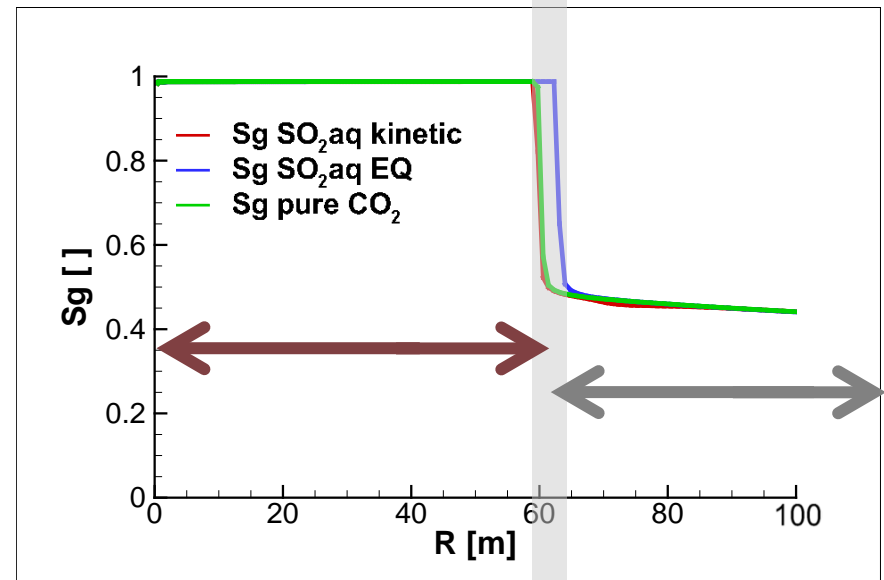


saturation

in the presence of water

→ all SO_2 dissolves

delay of EQ compared to kinetic



dry out zone

two phase
flow zone

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saturation

in the presence of water

→ all SO₂ dissolves

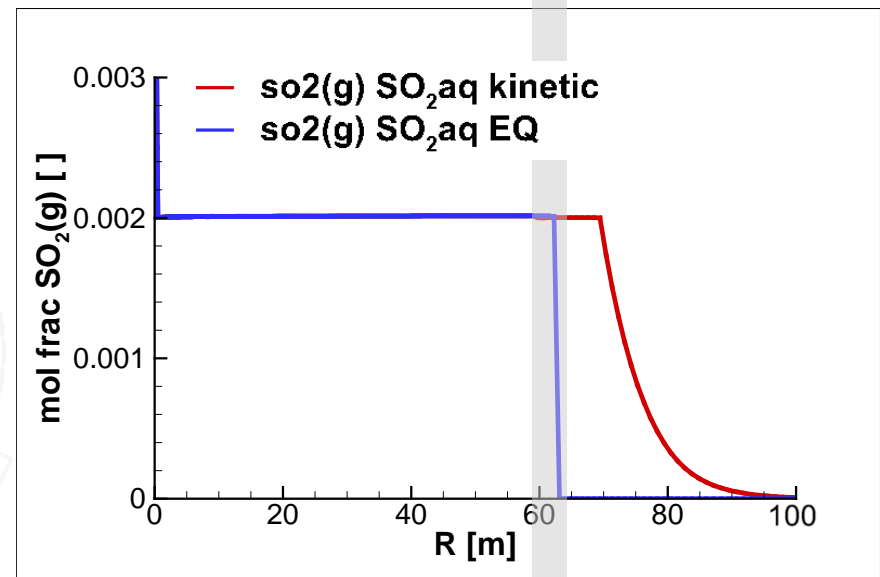
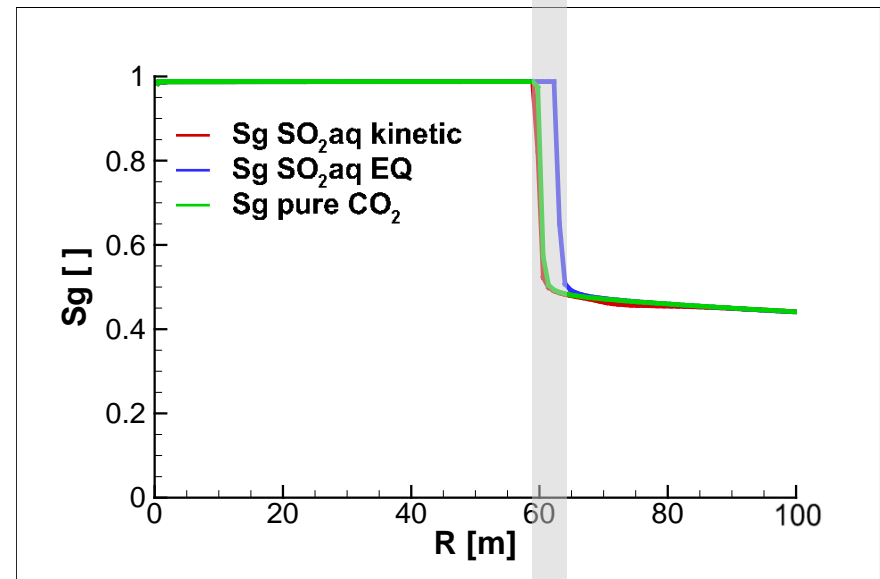
delay of EQ compared to kinetic

SO₂ in gas phase

EQ – contact of water

→ all SO₂ dissolves

kinetic – delay of SO₂ dissolution

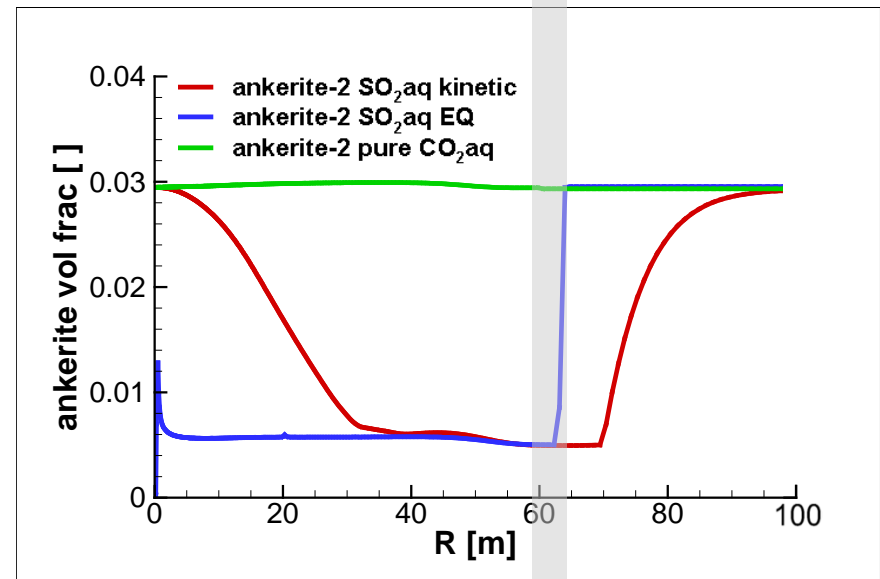




ankerite CaFe[CO₃]₂

EQ – difference to pure CO₂
til 60 m

kinetic – compared to EQ, first
smaller dissolution due
shorter residence time
→ diff. to pure CO₂ til 90 m



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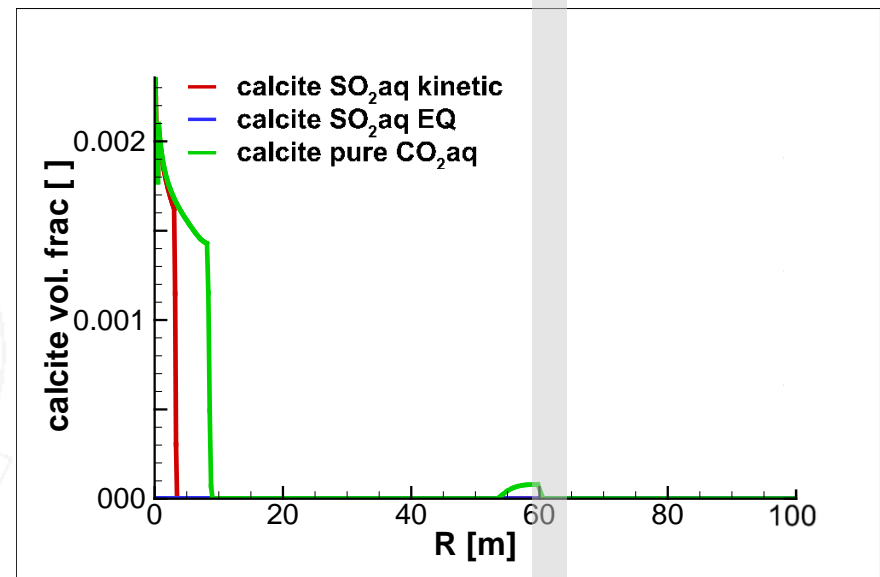
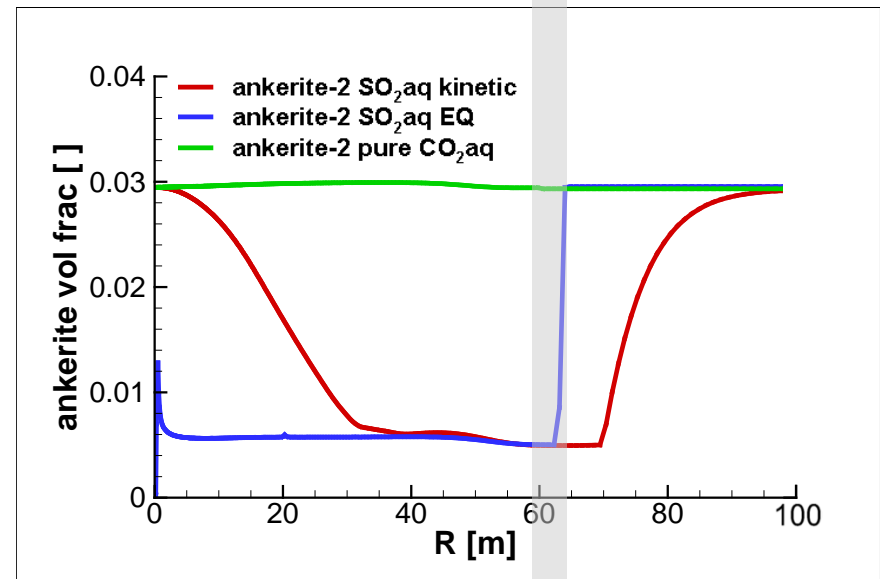


ankerite CaFe[CO₃]₂

EQ – difference to pure CO₂
til 60 m

kinetic – compared to EQ, first
smaller dissolution due
shorter residence time
→ diff. to pure CO₂ til 90 m

calcite CaCO₃
practically negligible

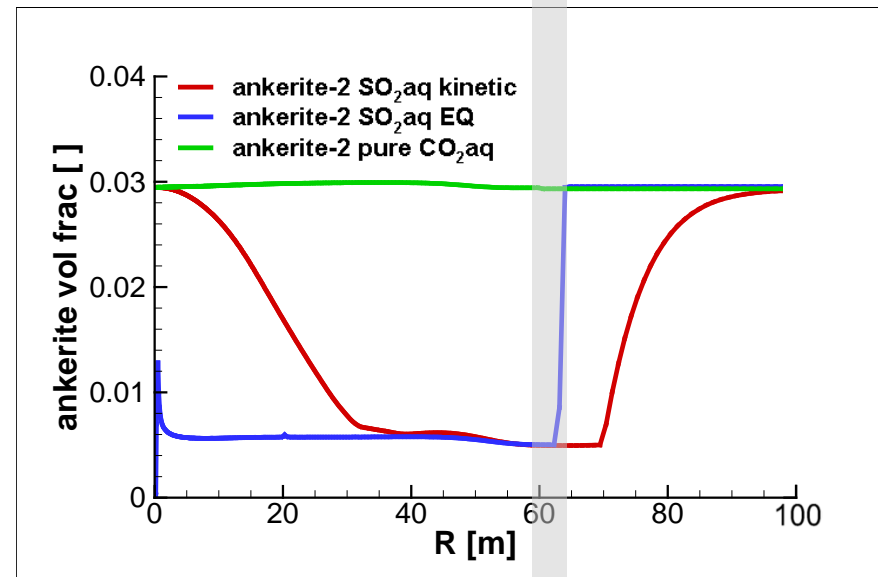




ankerite $\text{CaFe}[\text{CO}_3]_2$

EQ – difference to pure CO_2
til 60 m

kinetic – compared to EQ, first
smaller dissolution due
shorter residence time
→ diff. to pure CO_2 til 90 m

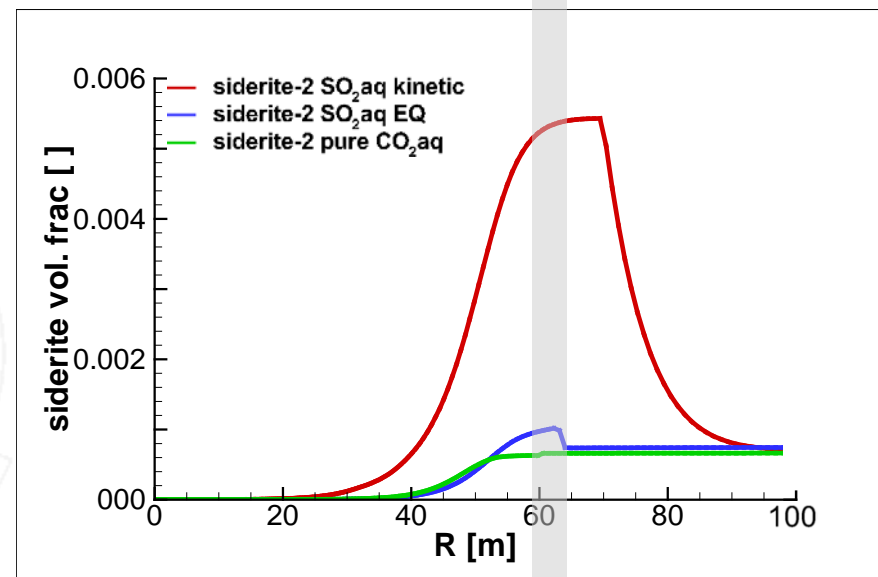


siderite FeCO_3

further away, small precipitation
using Fe released by ankerite

EQ – precipitation similar to pure
 CO_2 case

kinetic – higher precipitation

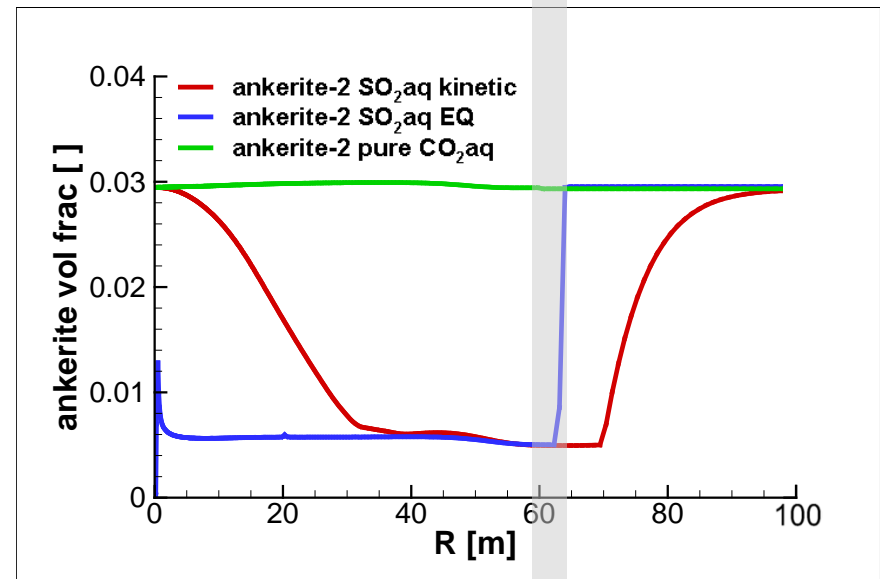




ankerite CaFe[CO₃]₂

EQ – difference to pure CO₂
til 60 m

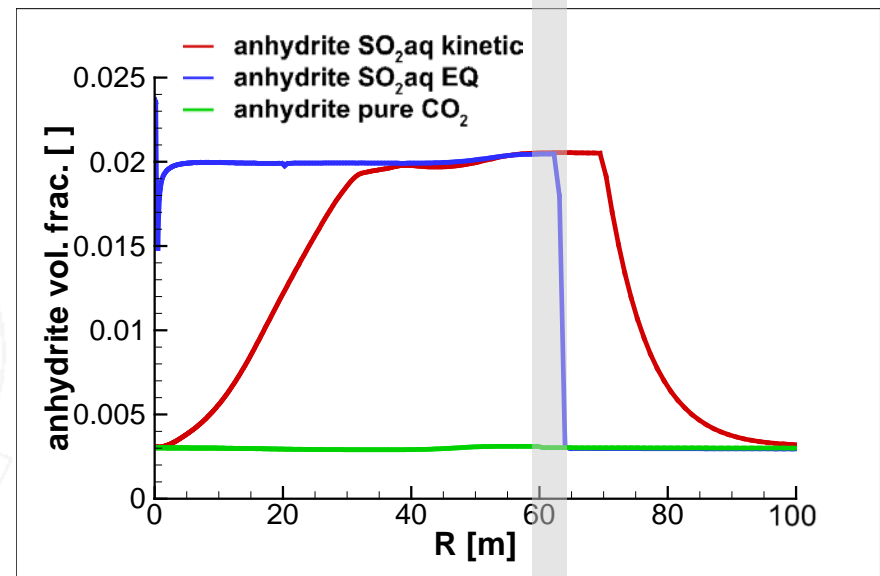
kinetic – compared to EQ, first
smaller dissolution due
shorter residence time
→ diff. to pure CO₂ til 90 m



anhydrite CaSO₄

– free Ca²⁺ from
ankerite CaFe[CO₃]₂
precipitates as CaSO₄

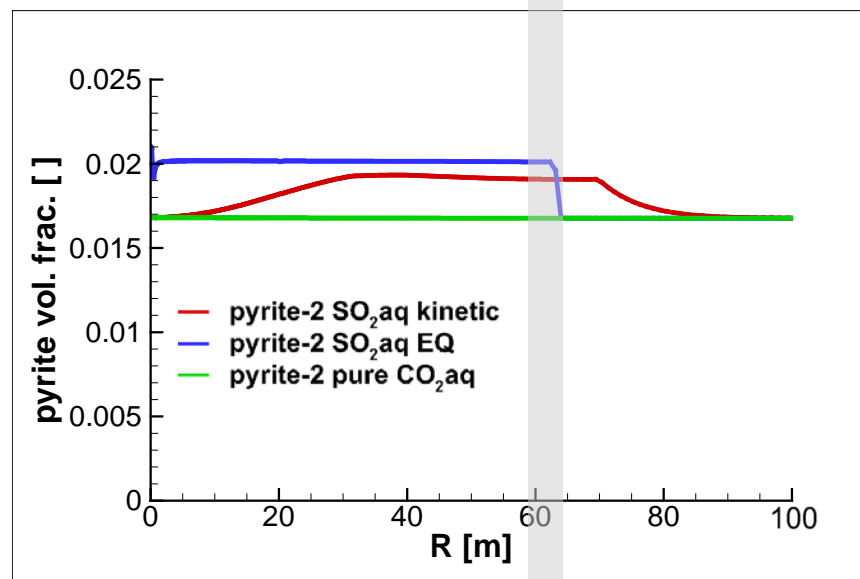
– delay in kinetic simulation
compared to EQ





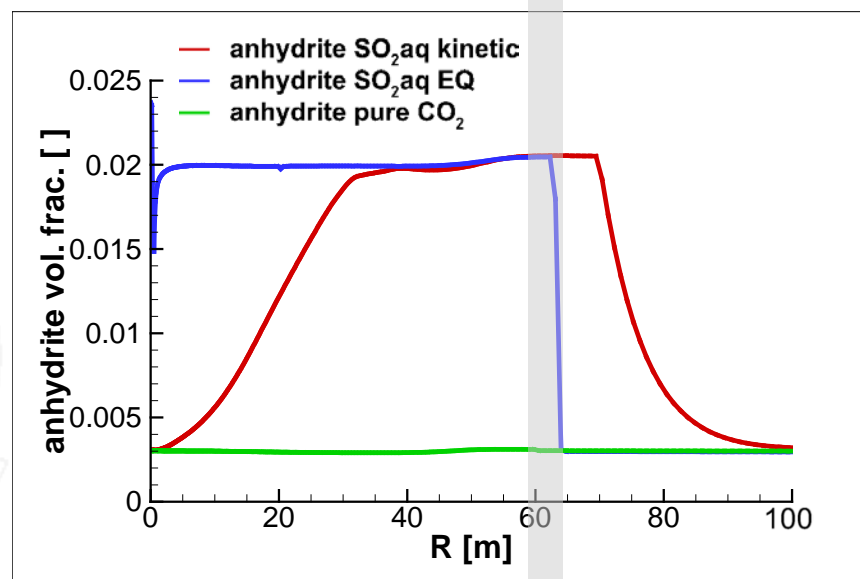
pyrite FeS₂

- precipitates parallel to anhydrite
- Fe from ankerite



anhydrite CaSO₄

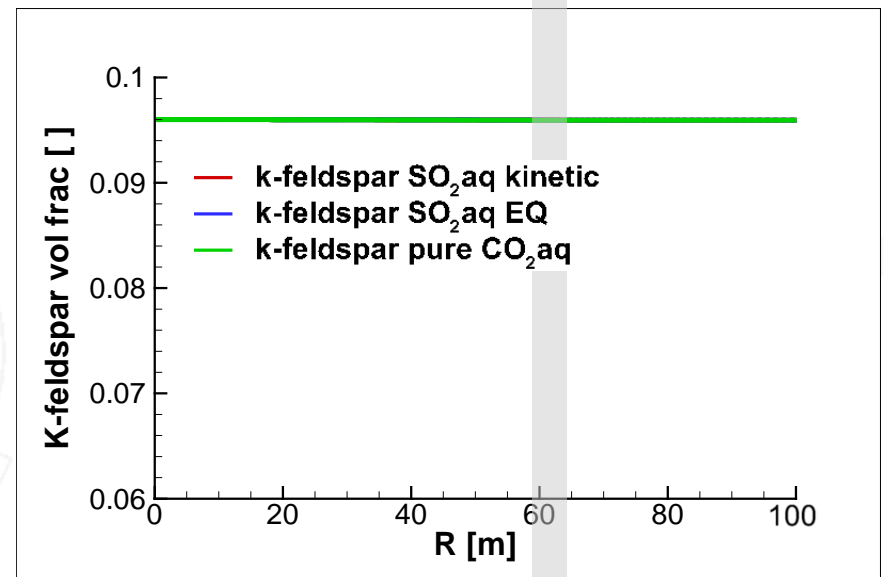
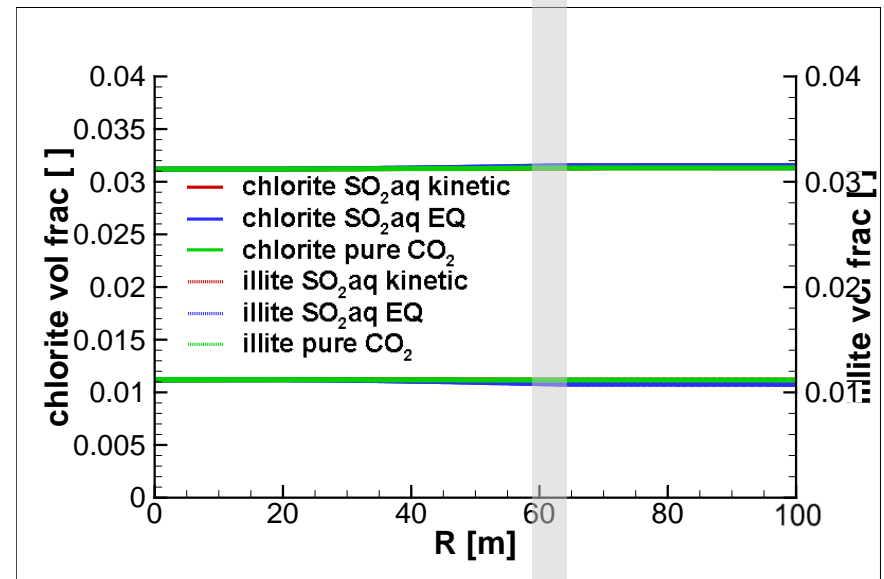
- free Ca²⁺ from ankerite CaFe[CO₃]₂ precipitates as CaSO₄
- delay in kinetic simulation compared to EQ





chlorite, $(\text{Mg,Fe})_3(\text{Si,Al})_4\text{O}_{10}$
 illite, e.g. $\text{K}_{0.85}\text{Al}_{2.85}\text{Si}_{3.15}\text{O}_{10}(\text{OH})_2$
 – no changes

K-feldspar, KAlSi_3O_8
 – no changes

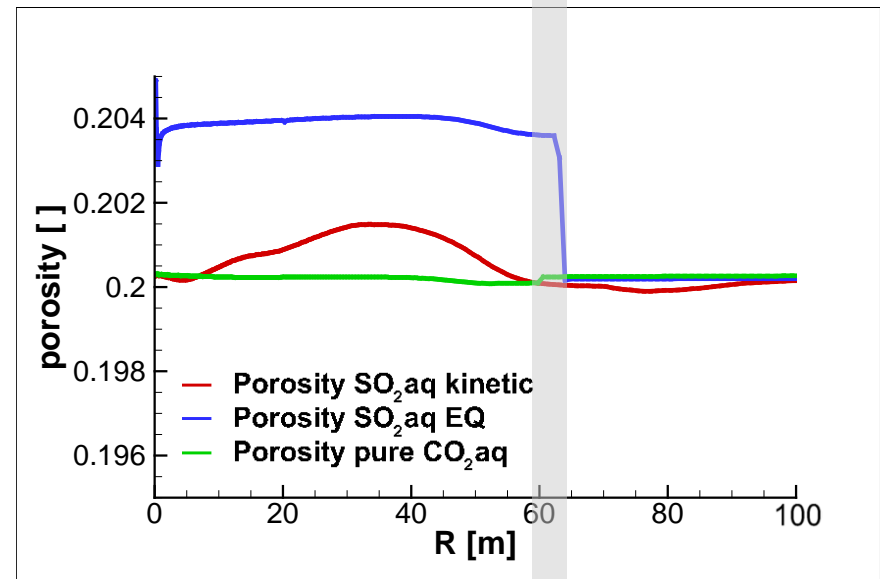


CO₂



porosity

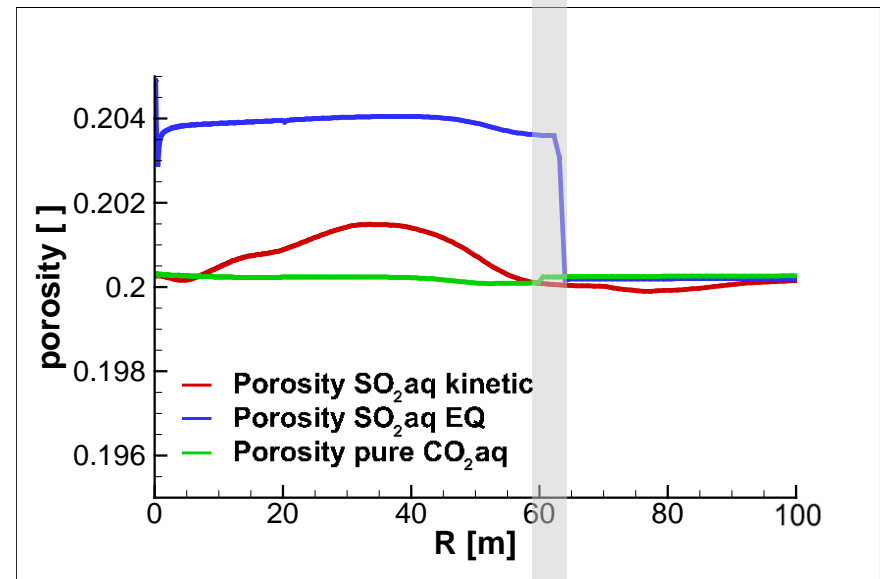
- determined through Ca : CO₃,
i.e. 1 : 2 as ankerite CaFe[CO₃]₂
more ankerite dissolves than
anhydrite CaSO₄ precipitates
→ increase in porosity





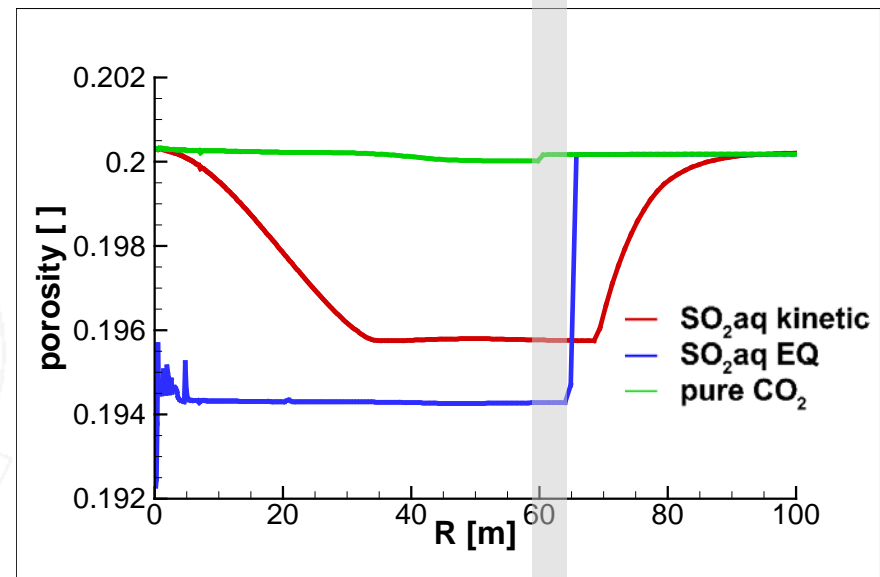
porosity I

- determined through $\text{Ca} : \text{CO}_3$, i.e. **1 : 2** as ankerite $\text{CaFe}[\text{CO}_3]_2$
more ankerite dissolves than anhydrite CaSO_4 precipitates
→ increase in porosity



porosity II

- with initial carbonate as calcite, i.e. **$\text{Ca} : \text{CO}_3$ as 1 : 1**
→ decrease in porosity as anhydrite has larger mole volume compared to calcite





conclusions

SO₂ rather small impact

results dependent on

- **real reservoir conditions,
esp. carbonates and porosity**
- **chosen model**

kinetics versus equilibrium conditions

choose your models wisely

outlook

further model validation
with Heletz data

further sensitivity analysis

influence of other impurities

1D → 2D reservoir model



J. Bensabat, EWRE

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disclaimer

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thank you

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